

An Investigation of 2013 M_w 7.7 Awaran Earthquake, Pakistan

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Abstract. Earthquake is a major disaster responsible for vast losses both life and infrastructure. Pakistan is one of the highly earthquake prone areas in Asia. The present study is carried out to investigate the dynamics of disastrous Awaran earthquake. The 2013 M_w 7.7 Awaran earthquake and its M_w 6.8 aftershock caused numerous loss of lives and vast infrastructure damages. The earthquake triggered along Hoshab fault rupturing 230 km of the fault. The rupture propagated at 3 Km/s on average and was concentrated in top 10 km with no major displacement in the underlying decollement. The rupture released a cumulative moment of 5.4×10^{20} Nm. This study reveals that although Hoshab fault originated as thrust fault, the slip was purely strike slip during the earthquake and that the rigid block rotation of southeastern Makran is responsible for the Awaran earthquake. The study concludes that the earthquake significantly increased the coulomb stress on Makran mega thrust and strike slip faults in Chaman fault system, hence increasing the risk of a major seismic event. Therefore, in order to prevent major loss of lives and infrastructure damages; designing of new building codes, reassessing the seismic hazard of the region and marking of hidden faults is of utmost importance.

1. Introduction

The Awaran earthquake with a magnitude of 7.7 M_w occurred on 24th September, 2013 [1] (Fig. 1 a,b). The depth of epicenter was 15 Km and the epicenter was located 61 Km North-North East of Awaran and 113 Km North-West of Bela [2]. The shaking lasted for about a minute creating panic in the region. The earthquake resulted in about 825 casualties and vast infrastructure damage (Fig. 2a). The earthquake was followed by intense sequence of aftershocks with a major 6.8 M_w aftershock on September 28th, 2013 which inflicted more damage to an already crisis hit area and atleast 22 people died as a result of this aftershock. The areas that were affected were Awaran, Tirtej, Gashkore, Nok Jo, Parwar, Dandar and Hohab [3]. About 33000 houses were damaged [3] and 300,000 people were affected as the result of earthquake [4]. The earthquake occurred on one of the southern strands of Chaman Fault System and was result of oblique-strike slip type motion [3]. The earthquake was felt all across Pakistan and in neighboring Afghanistan, Iran and India. Most of the population in this region resides in structures that are highly vulnerable to shaking caused by earthquake. Because of this powerful earthquake, a small island was created in the Arabian Sea which was mostly composed of mud and sand (Fig. 2b). The island was created 600 m offshore of the Gwadar coast, Southern Pakistan. The island was 200 m long, 20 m high and 100 m wide and was primarily composed of mud. The island was created because the violent shaking creates fractures and fissures in sediments which causes the pressurized methane gas to escape. Methane gas in-turn forces the overlying sediments to form a dome like structure which ultimately rises above the sea floor [5]. The earthquake occurred in one of the hazardous plate boundary area in the world. It is a zone of transition characterized by northward subduction of Arabian Plate beneath Eurasian Plate and also where Indian plate is sliding past Eurasian plate in northward direction.

Many major earthquakes have occurred in this region inflicting widespread damage (Table 1). The largest recorded event in the region is the M_w 8.1, 1945 Makran earthquake that generated a significant Tsunami. The event ruptured the subduction mega thrust. The second most significant event was M_w 7.5 Quetta earthquake that occurred on the Ghazaband fault. This study focuses on the review of Awaran earthquake dynamics.

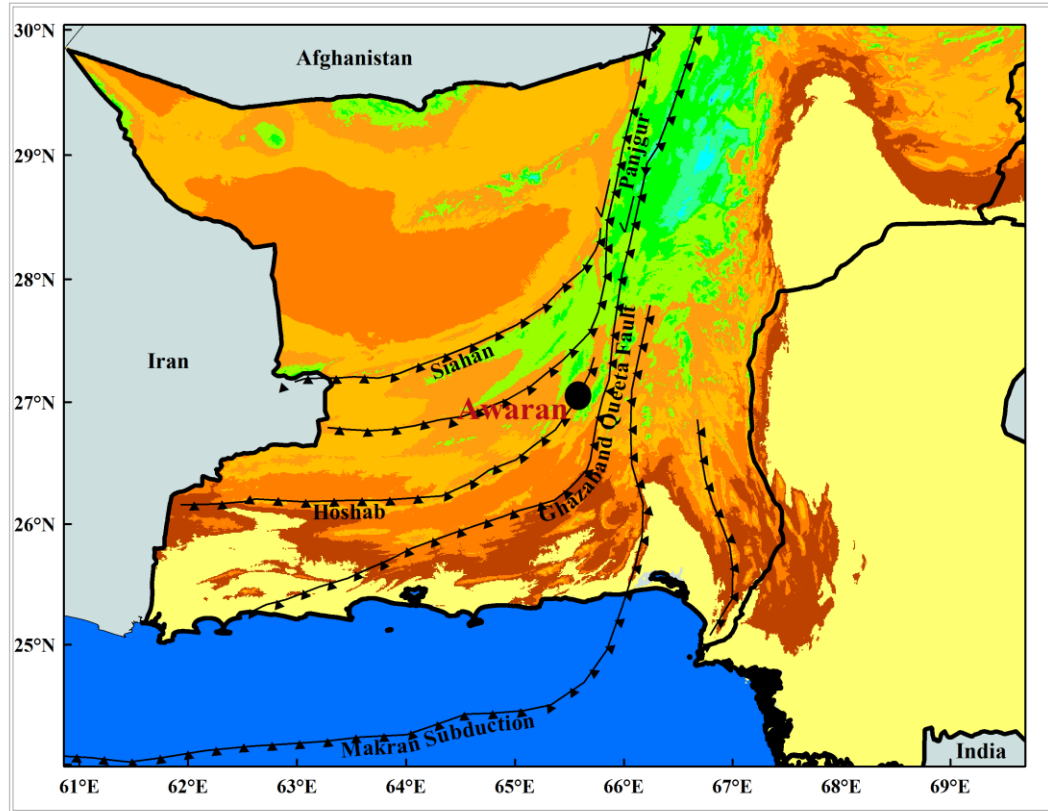


Figure 1a. Tectonic Map of Balochistan, Pakistan. Black circle represents location of Awaran Earthquake

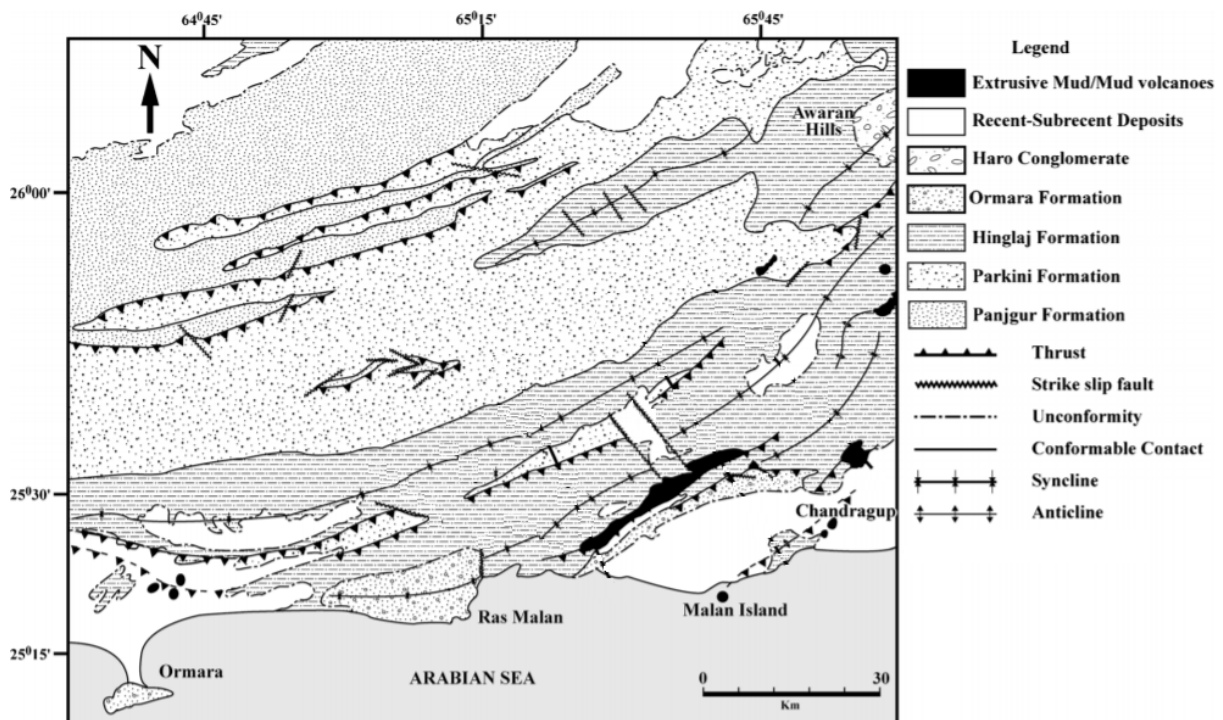


Figure 1b. Regional Geological Map of the area (Modified after Kassi et al. 2014) [6]

2. Tectonics of the Area

Pakistan is one of the most earthquake prone area in Asia and has high density of active faults. It is located in the intersection of three plate boundaries namely Indian, Eurasian and Arabian. In the southern region, subduction zone exists along Makran coast [1]. Makran plate boundary is a zone of wide deformation in southwestern Pakistan and southeastern Iran (Figure 1b). It extends for 400 Km NS and 1000 Km EW and accommodates a 3cm/yr northward subduction of Arabian and Ormara plates under Eurasian plate [7]. The boundaries of Makran Subduction Zone are mostly formed by transpressional strike-slip systems. Ornach-Nal fault system marks the eastern boundary whereas Minab-Zendan fault system marks the western boundary. Due to compression from Arabian plate, east-west directional structural trend is present in Makran zone and Chagai arc. It is subducting at a rate of 19.5 mm/yr with a subduction angle ranging from 2° to 8° [1]. Makran mega thrust was ruptured by a M_w 8.0 earthquake in 1945. Chaman fault system is a 1200 Km long, left-lateral fault system accommodating 3cm/yr northward motion of India with respect to Eurasia [8]. It extends south towards the subduction zone through Ghazaband and Ornach Nal fault systems [9, 10]. Chaman fault system is left-lateral transform plate boundary separating Indian and Eurasian plates. Chaman fault abruptly terminates the western edge of the Chaman fault system [9]. Chaman fault system comprises of Ornach Nal, Ghazaband and Chaman sinistral strike slip faults [11]. Southwards, Chaman fault connects with Siahan, Panjgur and Hoshab faults which are faults of Makran ranges and these faults accommodate; the shear between India and Eurasia and shortening due to convergence of Arabia and Eurasia [11]. Chaman and Ornach-Nal faults that are left lateral strike slip faults govern the colliding Indian and Eurasian Plate tectonics. Chaman fault comprises of four segments; Active fault, Chaman fault zone, Chaman fault system, Chaman Transform Boundary [1]. Tectonics and structural style of Ziarat area is governed by 1300 Km long Sulaiman lobe, Chaman fault and Sibi trough. Chaman fault, Kalat-Chinjin Structure, Harnai-Tarta Structure, Ghazaband Zhob Structure and Mach structure are local fault systems present around Quetta [1]. Seismic activity along coastal belt is concentrated on Ornach-Nal and Pub faults. A little south of Ormara and Pasni lies the 225 Km long Makran coast thrust fault [1]. According to Szeliga et al. 2012 [10], several faults in the region have shallow apparent locking depths (< 5 Km).

Table 1. Major earthquakes in Southern Pakistan

Location	Year	Magnitude	Casualties
Ali Jaan	1935	7.8	30,000 – 60,000
Arabian Sea	1945	8.1	4000
Quetta	2008	6.4	170-200
Dalbandin	2011	7.2	03
Awaran	2013	7.7	825
Awaran	2013	6.8	22

3. The Awaran Earthquake

The 2013 Awaran earthquake was a strike slip event that occurred on Hoshab fault [12]. Average slip was 4.2 m along the main trace of fault with an 11.4 m local maximum offset [12]. Maximum displacements of 16m were noted [13]. The earthquake occurred in south-eastern Balochistan and ruptured ~ 230 Km of the Hoshab fault, which is arcuate northwest-dipping fault [14]. Hoshab fault is located in Makran accretionary complex. 3 cm/yr of convergence between Arabia and Eurasia is being accommodated by Makran convergence zone [15]. The area is characterized by Panjgur, Naib Rub and Hoshab fault system (PHNFS) which are east west oriented reverse faults [13]. Fault geometries are well defined in earthquake region [16]. The Awaran earthquake was unusual in a sense that a curved reverse fault ruptured in mechanically weak

Makran Accretionary Prism. World View 2 high-resolution satellite imagery analysis also confirms a left-lateral surface rupture. Arcuate rupture corresponding to slip along the Hoshab fault was also confirmed using Landsat-8 imageries [13]. A portion of the accretionary prism where the slab is only 20 Km deep was also ruptured thereby suggesting that almost no crystalline crust ruptured in the earthquake. The earthquake ruptured a structurally segmented fault and involved a $\sim 6:1$ strike-slip to dip-slip ratio [14]. The rupture was associated to shallow depths from 0-7 Km. Most of energy (almost 80%) was released on south side of hypocenter whereas there is reduced slip in the north. Slower rupture propagation towards the north side is also confirmed by Teleseismic Inversion results [13]. The rupture released a cumulative moment of 5.4×10^{20} Nm hence corresponding to a M_w 7.75. It begins on a 70° westward dipping fault and then propagated to the 40° northwestward dipping fault segments to the southwest. The study confirms that during the seismic phase, measured slip occurred excluding the possibility of afterslip following the earthquake. The study also shows non-double couple moment. [17] (Fig. 3).

According to Jolivet et al. 2014 [8], the 2013 rupture was majorly left-lateral strike slip and involves a minor oblique reverse motion. The majority of slip occurred at shallow depths of $< \sim 10$ Km. Large $M_w > 7$ upper plate earthquakes have not been observed in the region before. The earthquake was triggered on a 45° dipping fault plane and depicts strike-slip movement. These faults, which are regional and moderate angle, initially originated as thrust faults but now depicts near horizontal slip [18]. According to Avouac et al. 2014 [17] and Barnhart et al. 2015 [14], there is a switch between thrust and strike slip faulting in the kinematics of Hoshab fault slip during successive earthquake cycles and that the fault ruptured in left lateral motion for 200+ Km and it dips between 47° & 73° to the northwest. Studies by Zhou et al. 2015 [12] suggests that the failure is not solely constrained to Hoshab fault but occurs in a distributed fashion thus providing an explanation to lack of seismicity previously and also to long repeat times. Younger and thicker sediments have more off-fault deformation [15]. Both reverse slip and strike slip deformation in rupture area was depicted by Quaternary geomorphic indicators. Arcuate geometry of Hoshab fault along with weak friction of Makran Accretionary Prism and time variable loading condition from plate interactions and adjacent earthquakes are responsible for the quaternary bimodal slip along the Hoshab fault [14]. Barnhart et al. 2015 [14] derived horizontal surface displacement fields by applying sub-pixel correlation to high-resolution optical imagery. The results showed that two structurally dissimilar sections of the Hoshab fault were ruptured due to the 2013 event. Geomorphic indicators express that the strike-slip fault trace geomorphology is pre-dominant along the northern sections, reverse-displacement along the southern section and reverse-type to strike-slip morphologies in center of rupture. The study expresses that variable slip with strike-slip motion, dip-slip motion or through going rupture may occur in individual sections of Hoshab fault as was the case in 2013 event. This study signifies that in the upper plate of Makran Subduction Zone bimodal slip might have accommodated strain portioning on a single fault. The 2013 surface rupture does not extend onto adjacent en-echelon faults but is rather restricted to single continuous trace [14].

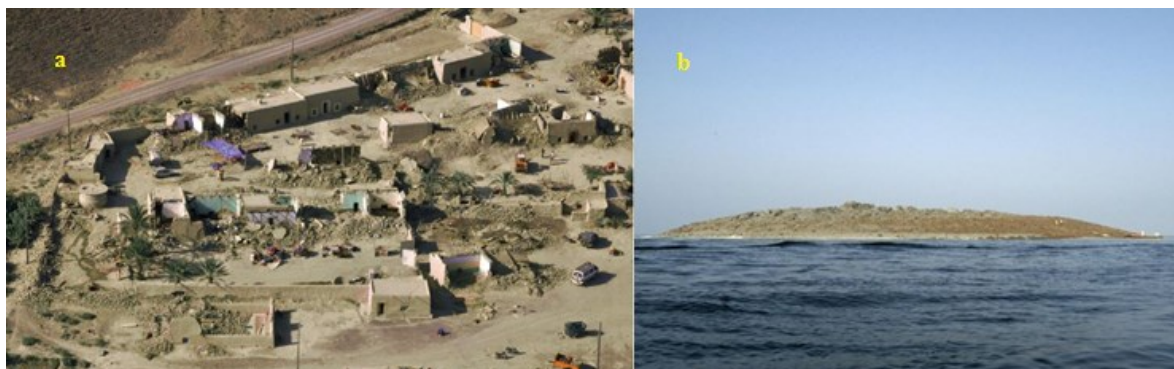


Figure 2. 2013 M_w 7.7 Awaran Earthquake. (a) Aerial view of destruction in Awaran area. (b) View of Island created due to Awaran earthquake [1]

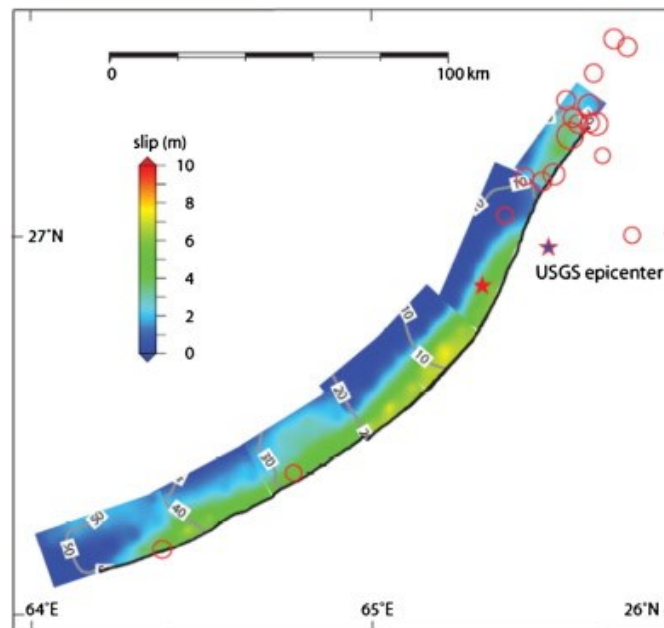


Figure 3. Finite source kinematic model showing average speed of rupture (3km/s) for this preferred model. Red circles show aftershocks with $m_b > 4$ over the first week following the main shock [17]

The rupture propagated at 3 Km/s on average and was unilateral. The fault trace suggests a westward bending of rupture with the arcuate tectonics of eastern Makran accretionary prism [17]. The 2013 earthquake indicate oblique strike-slip faulting as suggested by point-source focal mechanism results. Such type of strike slip faulting on a dipping structure is unprecedented. In addition, immense variation in the non-double component was revealed from different earthquake focal mechanisms (W-Phase, surface wave and body-wave solution). This suggests a significant source complexity [8]. Satellite studies suggest active thrusting accretionary prism including the Hoshab fault [19]. Despite the fact that long-term deformation is associated with active accretionary prisms, great earthquakes rarely occur there [8].

Hoshab thrust fault was reactivated with purely strike-slip motion in the event [17]. Total duration of rupture was 60 s with average velocity of 3 Km/s. The earthquake suggests that arcuate strike slip faults with $< 70^\circ$ dipping may go through strike slip rupture. Despite the fact that frictional strength is extraordinary weak in Makran accretionary prism, observations suggest that rigid block rotation of southeastern Makran, termed as “ball-and-socket” tectonic rotation is responsible for the 2013 earthquake [13]. It is also worth noting that > 7000 Km Chaman fault with a slip rate of more than 2 cm/yr has historically low level of seismicity [17]. Because of this earthquake; strike slip faults in the Chaman fault system, thrust faults of Kirthar range and portions of Makran mega thrust have come close for a near future failure. Studies by Avouac et al. 2014 [17] and Smith et al. 2013 [20] suggest that significant coulomb stress increased in the Makran mega thrust with significant increase in the deeper parts, which can slip aseismically. Coulomb stress also increased significantly on the thrust fault in the Kirthar range hence increasing the risk of a major seismic activity.

4. Conclusion

The 2013 Awaran earthquake that was triggered along Hoshab fault caused numerous losses of lives and infrastructure damage. The rupture propagated for more than 200 Km. The slip was purely strike slip during the earthquake. Most of the slip was confined to shallower depths of < 10 Km. Awaran earthquake is one of the most quantified slip event to occur on a curved dipping fault. Observations suggest that the 2013 earthquake was caused by rigid block rotation of southeastern Makran. As a result of this earthquake, the NS trending strike slip faults in the Chaman fault system are closer to being yield and also concise with the aftershocks that occurred in that

area. Therefore, there is a strong probability of a major seismic event in the Chaman fault system that can be devastating. It is very important; to reassess the seismic hazard of the region, designing of new building codes and marking of hidden faults, in order to prevent future loss of life and damages to infrastructure.

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